# 3D VSP – processing and integration with surface seismic

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## Introduction

Ever since the first 3D vertical seismic profile (VSP) acquisition was experimented in 1986 by AGIP, several 3D VSP volumes have been acquired [1-9]. Apart from the 3D VSP acquisition, a simultaneous VSP and 3D surface seismic acquisition has been attempted [10], and the results have been encouraging in terms of cost effectiveness and the enhanced imaging of the subsurface.

This work is an attempt to describe the simultaneous acquisition of 3D surface seismic data, the 3D VSP data and to detail the different steps that were followed to process the two data volumes. Attempts were made to enhance the bandwidth of the 3D seismic data using the VSP downgoing wavefield and integrate the 3D VSP volume with surface seismic to be able to interpret the two together.

## Acquisition of data

Simultaneous data acquisition of surface 3D and 3component 3D VSP data was carried out by Crestar Energy Ltd. (now Conoco), in October 2000 in Hanna area in Central Alberta. The geometry used for the 3D survey is the Mega Bin<sup>TM</sup> typical staggered brick pattern with a source spacing 70 m and a receiver spacing 140 m, sample rate 1 ms, record length 3 s and the energy source was dynamite. Using the same seismic source, and an 80 level (using Paulsson Geophysical tool), 3 component geophone cable in well 8-20, data was simultaneously recorded on the surface and the borehole. This acquisition represented approximately 3000 shots.

The 3D seismic programs (surface and VSP) were recorded to assist in defining sand presence and porosity development in the Lower Mannville interval (Ellerslie SS) found at a depth of between 1300.0 and 1350.0m. The 3D VSP recording was expected to (a) to help tie the seismic reflections to lithology and stratigraphic boundaries, (b) yield a high frequency image around the borehole (the fixed receiver array and its proximity to the reservoir are expected to improve image quality), and (c) help determine an improved subsurface velocity model.

#### Processing of 3D seismic data

The quality of the seismic was reasonably good. The processing sequence followed was pretty standard. A very effective QC used herein is the use of the Coherence Cube<sup>TM</sup> [11-16]. Coherence Cubes were generated at different stages in the processing flow and the subsurface features examined by slicing through the volume. Fig.1 shows a representative comparison of coherence slices depicting a channel running N-S through the slices. Notice that as the processing begins from the brute stack stage and goes through to residual statics application, the edges of the channel are better defined and similarly after migration the definition is still better.

## Processing of 3D VSP data

- (1) The processing of 3D VSP began with extraction of the z-component (vertical) from the 3-component recorded data. Next, the data was sorted in to different rings according to well-source distance around the well. The inter-ring distance was fixed at 50 m. This was continued for up to 32 rings or nearly 1600 m away from the well. However, on examining the sorted records, it was found that the amplitudes and signal-to-noise ratio was too low on the farther rings to be able to draw any meaningful conclusions from them. Fig.2 shows some representative records from different rings.
- (2) An obvious processing objective at this point is to extract the upgoing and downgoing wavefields and prevent the contamination of the data by the tube wave noise. Source statics corrections (derived from seismic) were now applied to the VSP gathers to bring them to a common datum. First breaks were picked automatically, but then corrected manually on each VSP gather. Maps for the first arrival times were prepared and QC'ed to estimate horizontal velocity changes and /or determine an anisotropy model. As part of the QC, these arrival times for different rings and channels were plotted. Fig.3 shows the arrival times for channel (depth level) 21 (550 m).
- (3) Zero offset VSP processing: The sources near the recording well were used to process a zero offset VSP which was used for initial calibration with the well logs and the surface seismic data. The raw zero-offset VSP data traces were edited and sorted by receiver depth. This was followed by wavefield separation, waveshaping deconvolution and corridor stack. The separation of the total wavefield into component wavefields (downgoing and upgoing P waves) was effected by an *optimisation inversion* method [17-19]. The methodology of using this approach contains several steps. As the first step, the different wavefield wavefield may contain noise (irregular) but also there may be some regular data wavetrains, which could be extracted from the total wavefield. In other words, the residual wavefield structure shows the quality of separation and adequacy of the model to real data.

Fig.4 shows the correlation between logs, upgoing VSP wavefield and surface seismic. Key formation tops are seen tracked across logs, real VSP (with the same depth index) the corridor stack (time) and seismic. Clearly, the frequency of the VSP is much higher than the seismic. Also, interspersed is bandpass filtered corridor stack to obtain a reasonable frequency match with the seismic. However, even for 5-40 bandpass, the seismic seems to be void of any reflection detail seen on the corridor stack around the zone of interest. With the known phase and polarity of the VSP trace, this correlation lends confidence to the surface seismic interpretation but, at the same time, also suggests that the frequency content of the seismic needs to be enhanced.

- (4) Offset VSP processing: Travel time inversion was used to create a layered anisotropic velocity model for each ring. To begin with, an isotropic layered velocity model was created from zero offset VSP first arrivals. This model consisted of 18 layers with the thicknesses ranging from 50 to 150 m. For different offsets, the first arrivals were calculated and overlayed on the real data. As seen in Fig.5, the isotropic velocity model derived from zero offset VSP first arrivals does not fit offset data. To get a better match, an anisotropic layered velocity model is used. For each ring, a vertical axis of anisotropy is considered with elliptical anisotropy associated with each layer. The elliptic anisotropy coefficients V<sub>H</sub>/V<sub>V</sub> (V<sub>H</sub> = horizontal velocity, and V<sub>V</sub> = vertical velocity) are determined by minimizing the difference between real first break time values and those of the model. The minimum of a misfit function was found by using a Monte Carlo approach (global optimisation) first and then local optimisation. After going through this exercise the average difference between the real and model arrivals is less than 1.5 ms. Fig.6 shows the actual wavefield and model first breaks for the anisotropic model. For different rings the anisotropic coefficients were found to be between 0.85 and 1.15 and in the shallow part rises to 1.25.
- (5) The velocity model so generated was used to calculate the upgoing and multiple downgoing travel times, to put them into each VSP gather and help during wavefield separation for each ring.
- (6) The total wavefield was separated into component wavefields, i.e. downgoing, upgoing, downgoing tubewaves and upgoing tubewaves.
- (7) VSP deconvolution (for both downgoing and upgoing wavefields). This helped in restoring frequencies of upgoing wavefield using the downgoing wavefield.
- (8) A close look at the downgoing and the upgoing wavefields, revealed an interesting observation. The wavelets embedded in these data do not appear to be uniform. This set us thinking about the causes of the non-uniformity in the wavelets. We plotted the shot point depth and found that there is a significant oblique N-S trend in the depth pattern and is not uniform. To try and equalize the embedded source wavelets, a matched filtering was done between a model wavelet and the embedded source wavelets in the data. This enabled us to have a globally consistent wavelet in the data. The model wavelet was chosen as the average source wavelet in the consistent zones.
- (9) VSP-CDP time-to-depth transformation for every ring was done to obtain vertical VSP profiles. For this, the velocity model determined in step 4 was used. Next, all depth images were stacked after making automatic time shift determination. This prevents any loss of higher frequencies and a crisp image was obtained for different offsets.
- (10) Stacking of ring images (with time shifts) to obtain a 3D VSP volume, still in depth.
- (11) Depth to time transformation and integrating with 3D VSP volume. Some vertical planes cutting through the 3D VSP time volume may be seen in Fig.7. This volume was then merged with the seismic volume

#### Conclusions

The processing of 3D VSP was an enriching experience, though quite labour intensive and time consuming. The vertical profiles through the 3D VSP volume exhibited higher resolution and so detailed and focussed reflections, especially at the target level. The deployment of a large number of borehole seismic receivers was expected to yield a larger reflection coverage around the borehole and seems to have paid off here. The results of this project have demonstrated the benefit of the integrated approach pursued which seeks to optimise the set exploration objectives.

The experience gained through processing, integrating the two data sets recorded simultaneously and evaluating the final results permit us to recommend that such simultaneous surveys conducted at an incremental cost (so commercially feasible) could provide detailed images sought around the borehole.

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After brute stack 984 n





Fig.1: Coherence time slices at different stages of processing



Fig.2: Tube waves as seen on records in different rings



Fig.3 : Planar view of first arrival times for channel 21



Fig5: First time arrivals for ring 21, offset 1050 m



Fig6 : First time arrivals for ring 21, offset 1050 m



Fig.4: Correlation of stratigraphy, VSP upgoing wavefield, well logs, VSP corridor stack and seismic



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